Contribution of water and tillage erosion to bright patches formation on the base of erosion modelling (Case study Trnavská pahorkatina Hill Land, Slovakia) Preliminary results

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Introduction

The relief of agricultural land is influenced by various geomorphic processes variable in space, time and intensity. Water and wind erosion are considered to be dominant, but the recognition of the influence of tillage erosion on soil redistribution (Bac 1928, Lobotka 1958, Lazúr 2001, 2005, cf. Van Oost et al. 2006) and landform evolution (e.g. Janicki et al. 2002, Van Oost et al. 2005, Stankoviansky 2008) is growing. The processes results in the variation in soil properties, uncovering of maternity rock and consequently to relief lowering. The areas with loess substratum visible on the surface have a great extent in the loessic areas in Slovakia. So-called bright patches are easily distinguishable due to contrast between bright loess and surrounding darker soils. So far mainly their occurrence, distribution and soil properties has been studied (e.g. Sobocká, 2002, Ilavská, Jambor 2005). Several suggestion on their origin have been proposed. Possible relationships with archaeological sites occurrence (Kohan 1993), relict (Linkeš et al. 1992) or recent water erosion (Fulajtár, Janský 2001) as well as polygenetic origin with contribution of eolian processes (Stankoviansky 1993) were presumed. However possible contribution of tillage erosion has been suggested recently (Smetanová 2008) more detailed research on bright patches genesis and relief development of loessic hill land is needed. The objective of the paper is to contribute to this discussion by study of

bright patches occurrence and modelling of erosion processes.

Methods

Two study areas (64 and 82 hectares) are situated in SW of Trnavská pahorkatina Hill Land in Danube Lowland. They represent typical relief of this part of hill land, which is divided by neoteconic fault activity into softly modelled depression and elevated plateaus connected with steeper slopes. There in the steepest parts (with average slope 2° and maximum slope 8.6°) the first study area is situated. In the west it borders with the riverbed of brook Ronava and in its northern part shallow dell occurs. The second area represent moderate relief of the plateau with average slope 0.65° (max. 2.6°). Chernozems on loess strongly affected by erosion are predominant and bright patches occur.

Bright patches were identified using visual interpretation of orthophotomaps and ortorectified aerial images. Detailed DEMs were created on the basis of topographic maps (1:10,000, 0.5–1m distance between contour lines) and GPS terrain measurements. The WATEM model (Van Oost et al. 2000) was used to asses water and tillage erosion rates and their spatial distribution. Values of rainfall erosivity and soil erodibility factor for water erosion modelling were previously published (Ilavská 2005, Ilavská,

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Jambor 2005, Styk, Pálka 2005). Crop erosivity factor was calculated as weighted three year average of the erosivity of particular crops and the area covered by them on the basis of available crop rotation data.

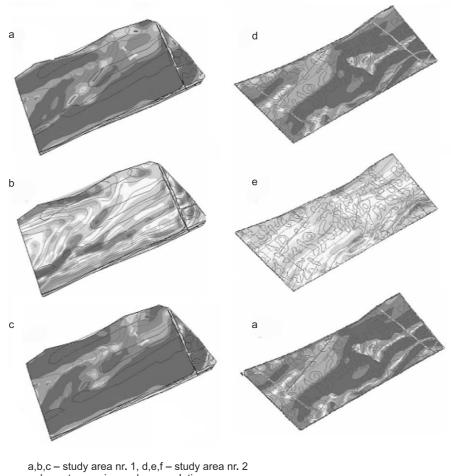
Tillage erosion is a diffusion-type process controlled by the change in slope gradient and intensity dependent on the value of tillage transport coefficient (Van Oost et al. 2000). Previously published values (Van Muysen at al. 2000) refering to applied mouldboard, tillage depth and average speed were used. The results of water erosion, tillage erosion modelling and their cumulative effect were compared with bright patches distribution.

Results and conclusion

The bright patches cover significant part of the studied areas (21.3% and 42.9% respectively). Long and narrow patches are predominant on the steeper slopes. Smaller in slope direction elongated patches prevail in the second study area. The highest water

erosion rates (with maximum 6.12 t ha⁻¹·a⁻¹) occur in the steepest lower parts of slopes in the first study area (Fig. 1a). Accumulation increases up to 6, 87 $t \cdot ha^{-1} \cdot a^{-1}$ there (32.8 $t \cdot ha^{-1} \cdot a^{-1}$ in the area close to riverbed of creek Ronava). The greatest rates of tillage erosion (5.51 t \cdot ha⁻¹ · a⁻¹) are connected with ridges and terrain edges (Fig. 1b), less intensive process can be observed close to field boundaries. Accumulation (maximum 3.32 t·ha⁻¹·a⁻¹) occurs in concave parts. Comparing the areas of erosion and bright patches occurrence their connection to the areas with highest tillage erosion rates is noticeable. They extent partly in the areas with tillage accumulation, where water erosion is more intensive and therefore antagonistic effect of both processes is responsible for the formation of bright patches. (Fig. 1c).

In the second study area both processes are less intensive (Fig. 1d, e) with the maximum water (tillage) erosion 2.4 (0.63) t \cdot ha⁻¹·a⁻¹ and accumulation 35.7 (0.95) t \cdot ha⁻¹·a⁻¹ rates. Bright patches are distributed within the whole slope without significant linkage to tillage erosion. They occur more often in areas



a,d – water erosion and accumulation b,e – tillage erosion and accumulation

c,f, - cumulative effect of water and tillage erosion and accumulation

Fig. 1. Water and tillage erosion and their cumulative effect

with higher water erosion or with acumulation. The cumulative effect of both processes is more dependent on water erosion (Fig. 1f).

Modelled cumulative rates (with maximum 6.68 and 2.3 t·ha⁻¹·a⁻¹) are lower than values estimated by Linkeš et al. (1992) by ¹³⁷Cs measurements in soil catena (study site near Voderady). The difference between modelled and measured values of erosion indicated the need for further evaluation of current erosion pattern. The evaluation of historical erosion and its variability is crucial for the study of bright patches formation. Further research is also needed to distinguish the role of particular processes, including tillage erosion which probably more influences the formation of long and narrow patches linked to terrain edges, ridges and tops.

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