

## Genetic interpretation of micromorphological features of gully loess-soil deposits (case study: Kolonia Celejów, E Poland)

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**Abstract:** Five complex soil profiles in gully system on Nałęczów Plateau (E Poland) were examined. They represent the loess sediments modified by soil processes and postpedogenic transformations. Two examined profiles represent natural and mature Luvisols with diagnostic illuvial horizon (Bt argillic) developed on the edge of plateau geomorphologic level and in the relict valley bottom. The other soil profiles reflect postpedogenic soil degradation, redeposition and secondary accumulation (deluvia, colluvia and/or proluvia). Record of these processes are the micromorphological features created by lithological processes (=primary loess and secondary pedoliths) or by pedogenesis (mainly bioturbation, illuviation, de- and recalcification). The important group of microfeatures are pedorelicts which clearly confirm genetic dependences of pedoliths (deluvia and proluvia) and soils developed *in situ* and located in higher hypsometric levels in the catchment.

**Keywords:** gully erosion, loess, pedogenesis, soil degradation, micromorphology

### Introduction

Micromorphological method is mainly used in pedology to describe the soil processes and specify the typology of soils (e.g. Bullock et al. 1985; Stoops 2003). Therefore, its main role is evidence the individual and specific microforms created by definite soil process or the group of microfeatures which reflect the multi-stage evolution of soil cover. For this reason, micromorphology has found wide application in studies of polygenetic sediments and soils, in particular the Quaternary loess-paleosol sequences (e.g. Kemp 1999, 2001, Mroczek 2008).

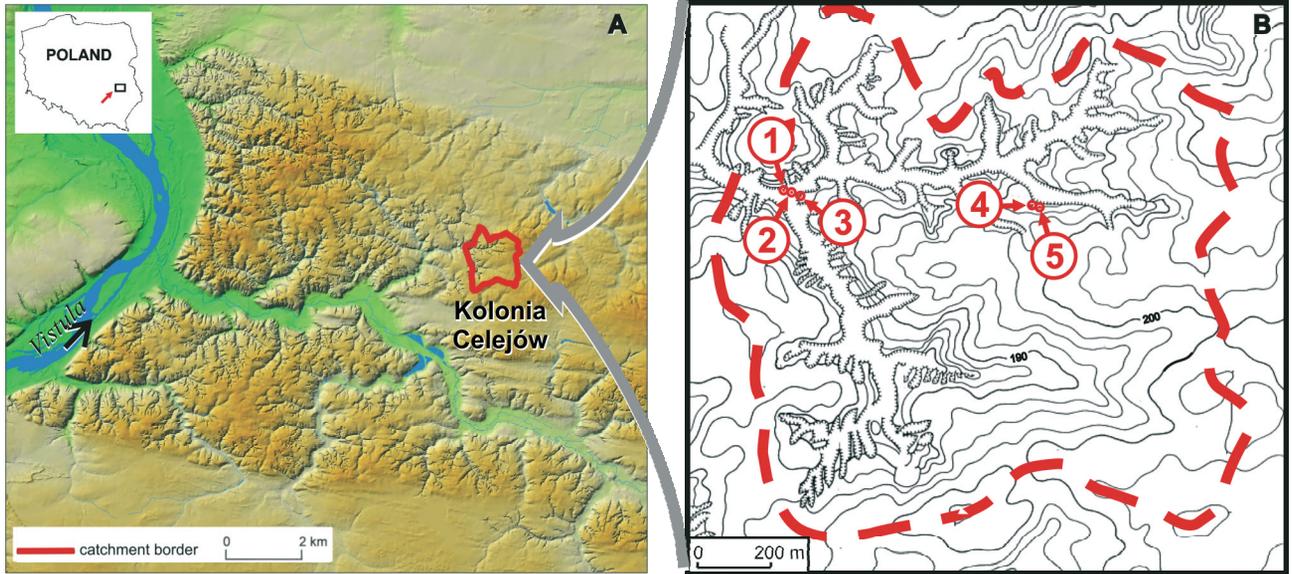
Usually, very complex nature – like described above – have gully sediments filling the bottoms, slopes and fossil piping channels or forming the accumulation layers next to the erosion edges. Their primary location is commonly associated with degraded and redeposited layers originally situated in higher geomorphologic areas in the catchment. Finally, their homogeneity impedes the indication of primary sources of gully sediments, especially in clastic materials filling the bottoms as so-called *proluvia*. However, the use of micromorphology in research of redeposited sediments enable to indicate

the alimentation sources through the identification of specific older and relict microfeatures created by litho- and pedogenic processes. These microforms allowed to detail the genesis and specify the ways and character of transport (e.g. Mroczek 2008).

### Materials and methods

The selection of the study areas was based on previous field works and on earlier studies in the western part of northwest part of Lublin Upland (e.g. Rodzik & Zglobicki 2000; Rodzik et al. 2009). The area is located in Nałęczów Plateau, adjacent to the Vistula River valley (Fig. 1). The main surface rocks are loess and loess-like sediments (max. thickness 30 m), where the relative relief reaches up to 90 m and the average density of the gully network – ca. 2.5 km km<sup>-2</sup>.

Micromorphological analyses were conducted in four loess-soil profiles located within the gullies dissecting the system of dry valleys in the Kolonia Celejów on the Nałęczów Plateau (Rodzik et al. 2009). The examined exposures differ mostly in their geomorphological situation, degree of soil develop-



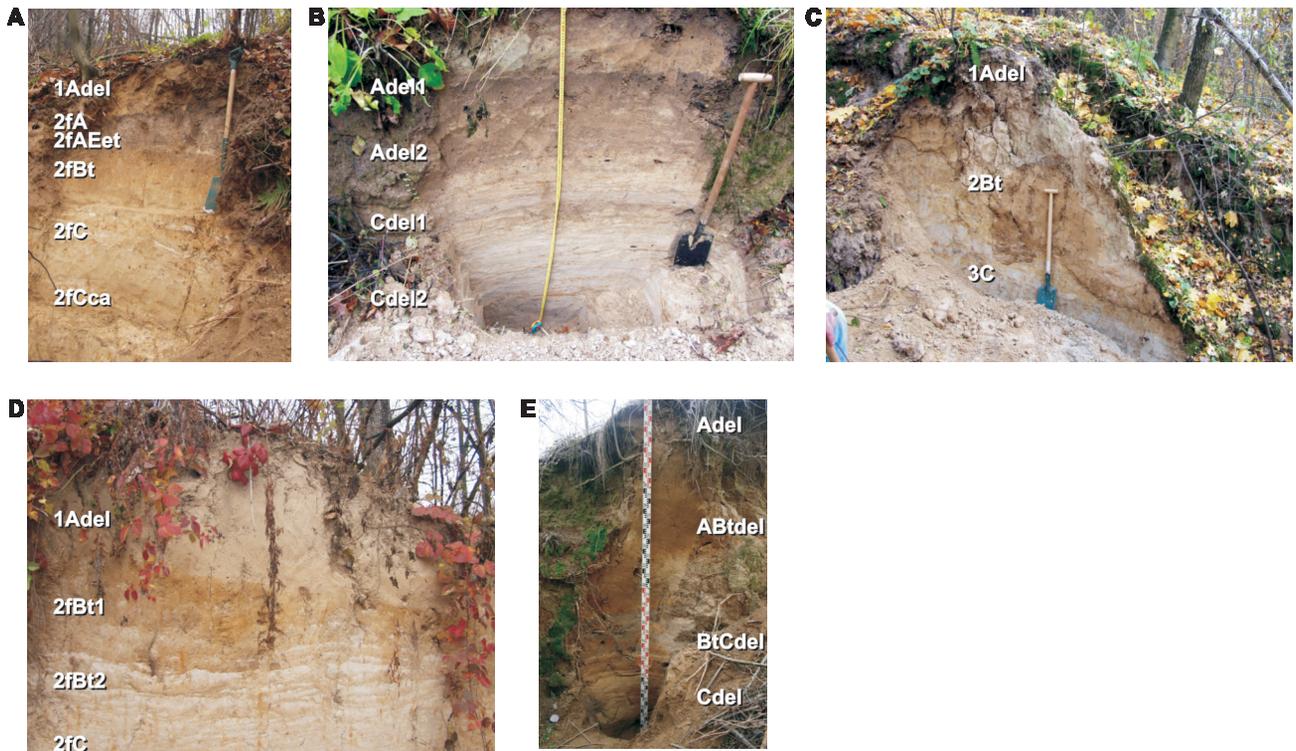
**Fig. 1.** Location of the Kolonia Celejów on western part of Nałęczów Plateau (A) and soil profiles in the gully, mentioned in the study (B)

ment, and kind of possible postpedogenetic transformations (Fig. 2). Soil horizons were described using the terms outlined by Reuter (2000). Soil profiles description is as follows:

- profile 1: Buried Luvisol with the profile 1Adel-2fA-2fAEet-2fBt-2fC-2fCca, developed in the top of loess remnant (deluvial loess) in the escarpment of gully dissecting the valley bottom;
- profile 2: Initial deluvial soil with the profile Adel1-Cdel1-Cdel2, developed on the gully bot-

tom and accessible in a dissection of proluvial deposit cover;

- profile 3: Tripartite deluvial-colluvial sequence with the profile 1Adel-2Bt-3C, accessible in the gully escarpment, in the inactive suffusion kettle dissecting the valley slope;
- profile 4: Truncated and buried Luvisol with the profile 1Adel-2fBt1-2fBt2-2fC-2fCca, accessible in the gully escarpment, in the reactivated suffusion kettle dissecting the valley slope;



**Fig. 2.** Examined soil profiles: A – buried Luvisol in the escarpment of gully, B – initial deluvial soil in gully bottom, C – tripartite deluvial-colluvial sequence in the gully escarpment, D – truncated and buried Luvisol in the reactivated suffusion kettle, E – sequence of soil deluvia in the reactivated suffusion kettle

- profile 5: Sequence of soil deluvia with the profile Adel-ABtdel-BtCdel-Cdel filling the reactivated suffusion kettle (opposite the profile 4).

The undisturbed samples were collected using standard (80×60×40 mm) Kubiena tins for safe transport back to laboratory for impregnation with crystic resin. The technique adopted the approach of Lee & Kemp (1992) and Mroczek (2008) to produce approximately 30-micron thick thin sections (7×5 cm<sup>2</sup>). More than 30 thin sections were analysed under plane and polarized light. Samples were described using the terms outlined by Bullock et al. (1985) and Stoops (2003).

## Results and conclusions

Micromorphological analyses (see Table 1) enable us to draw the following genetic conclusions:

1. The mature soils (profiles 1 and 4) belonging to Brownearths type contain preserved and rather well developed Bt horizon with features of argic diagnostic endopedon. Illuvial clay coatings and infillings in channels and pore space are its representative and diagnostic microfeatures;
2. The products of redeposition of soil horizons occur as the infilling of suffusion kettle (profile 5) and as deluvia (so-called agricultural diamictons, profiles 3, 4 and 5). This fact is mainly evidenced by pedorelict features of these deposits (according to Fedoroff & Goldberg 1982 and *vide* classification proposed by Mroczek 2008). From among all distinguished micromorphological features, the following belong to pedorelicts: illuvial papules and angular Fe and Fe-Mn nodules, and similarly looking amorphous pieces of charcoal. Their primary soil origin is confirmed by optical features and size resembling the forms occurring *in situ*, as well as pebble-like shape indicating rounding of primary microforms during their redeposition (rather short transport);
3. The Bt horizon of mature soil (profile 3) was redeposited *en bloc* without additional deforma-

tion. Micromorphological record confirms its illuvial origin and lack of connection with the overlying epipedon 1Adel (diamicton) and underlying endopedon 3C (primary loess). This horizon can be recognized as typical colluvia composed exclusively of material from the Bt-argic soil horizon.

Proluvia covering the modern bottom of gully (profile 2) are the products of redeposition of primary loess and soils developed on it, which were transported at longer distance and more mixed (homogenized) than the materials mentioned above. This fact is unambiguously indicated by pedorelict features, which are considerably more rarely found and individual microforms have smaller diameters, as well as by the occurrence of lithogenic carbonate microforms (microsparite and sparite crystals).

The conducted analyses enable us to evidence the record of natural postglacial pedogenesis in the form of Luvisol profiles with the diagnostic Bt-argic soil horizon in plateau-slope and valley bottom positions. Their age is older than that of gully system. The complex of micromorphological features distinguished in the C horizon (parent material) of these soils has remarkably lithogenic nature so they developed on primary loess.

Pedorelict features of the redeposited sediments indicate their alimentation sources. Moreover, they evidence a gradual obliteration of pedogenic features with the transport length. These features enable us to reconstruct the sequence of events: sedimentation → pedogenesis → degradation → transport → deposition, with partially still preserved their soil nature.

## Acknowledgements

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## References

- Bullock S., Fedoroff N., Jongerius A., Stoops G. & Turisna T., 1985. *Handbook for Soil Thin Section Description*, Waine Research Publ., Wolverhampton, England.
- Fedoroff N. & Goldberg P., 1982. Comparative micromorphology of the two late Pleistocene paleosols (in the Paris Basin). *Catena* 9: 227–251.
- Lee J. & Kemp R.A., 1992. *Thin section of unconsolidated sediments and soils: a recipe*. Thin Section Laboratory, Sediment Analysis Suite, Geography Department, Royal Holloway, University of London, Egham.
- Kemp R.A., 1999. Micromorphology of loess-paleosol sequences: a record of paleoenvironmental change. *Catena* 35: 179–196.
- Kemp R.A., 2001. Pedogenic modification of loess: significance for palaeoclimatic reconstructions. *Earth-Science Reviews* 54: 145–156.
- Mroczek P., 2008. *Interpretacja paleogeograficzna cech mikromorfologicznych neoplejstoceńskich sekwencji lessowo-glebowych*. Wydawnictwo UMCS, Lublin: 131 pp.
- Reuter G., 2000. A logical system of paleopedological terms. *Catena* 41: 93–109.
- Rodzik J., Furtak T. & Zgłobicki W., 2009. The impact of snowmelt and heavy rainfall runoff on erosion rates in a gully system, Lublin Upland, Poland. *Earth Surf. Proces. Landforms* 34: 1938–1950.
- Rodzik J. & Zgłobicki W., 2000. Wpływ układu pól na rozwój wąwozu lessowego. In: Radwan, S. & Lorkiewicz, Z. (eds.) *Problemy ochrony i użytkowania obszarów wiejskich o dużych walorach przyrodniczych*. Wydawnictwo UMCS, Lublin: 257–261.
- Stoops G., 2003. *Guidelines for Analysis and Description of Soil and Regolith Thin Sections*. Published by Soil Science Society of America. Madison, Wisconsin, USA: 176 pp.